Natronambulite from the Dohnoiri mine, Kanuma City, Tochigi Prefecture, Japan

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Abstract Natronambulite from the Dohnoiri mine was found as the second occurrence in Japan. It is an essential constituent of bedded manganese ores formed by very weak metamorphism. The other manganese minerals are braunite, rhodochrosite and rhodonite. They occur as minute grain aggregates forming fine bands with or without quartz. After the calculation of Li₂O and H₂O contents, EDS analyses are idealized as (Na,Li,Ca)(Mn,Ca,Mg)₄[Si₅O₁₄(OH,O)], where Na>Li, Mn>Ca, Mg and OH>O, substantiating the original interpretation of essential nature of Ca substituting Mn. Natronambulite and nambulite in Japan have been found in weakly to moderately metamorphosed bedded manganese ore deposits which are principally developed in such a siliceous rock as chert in Jurassic mélange. The ore minerals therefrom include braunite without exception. It is thus very likely that the oxidizing condition might have helped the concentration of lithium. Two natronambulites in Japan were found from a member of low grade siliceous bedded manganese ores in all the known Japanese localities. These two lithium-manganese silicates might have been confused due to the color with rhodonite or pyroxmangite, especially high iron materials of the former, suggesting the more common occurrences.

Key words: natronambulite, Dohnoiri mine, manganese ore.

Introduction

Natronambulite, (Na,Li)(Mn,Ca)₄[Si₅O₁₄OH], was first described from the metamorphosed bedded manganese ore deposit of the Tanohata mine, Iwate Prefecture, Japan by Matsubara et al. (1985), as the Na-dominant member of nambulite, (Li, Na)Mn₄[Si₅O₁₄OH], in which Li>Na (Yoshii et al., 1972). However, Matsubara et al. (1985) considered the mineral not to be a mere Na-dominant member but pointed out the essential nature of Ca replacing Mn. Prior to the first find of this mineral, the similar material had been found in the Kombat mine, Namibia, by von Knorring et al. (1978), who handled it as nambulite due to the reason that the original nambulite was chemically characterized as LiNaMn₈- $[Si_5O_{14}OH]_2$ with Li: Na=0.51: 0.49 (Yoshii et *al.*, 1972), while the Kombat material has Li : Na=0.40:0.60, and the authors considered the chemical difference between two to be too insignificant to create a new species.

Synthesis of LiMn₄ [Si₅O₁₄OH] was succeeded by Ito (1972) prior to the discovery of nambulite. He also acquired synthetic (Na,Li)Mn₄-[Si₅O₁₄OH], but failed to obtain synthetic pure Na end member, or NaMn₄[Si₅O₁₄OH]. According to his personal communication he remarked very slow reaction speed of gel mixture of employed ingredients for NaMn₄ [Si₅O₁₄OH] and possibility of disintegration of product into two, synthetic serandite (NaMn₂[Si₃O₈OH]) and pyroxmangite (Mn₇[Si₇O₂₁]).

The present natronambulite contains essential Ca in place of Mn as seen in the original analysis, substantiating the interpretation as to the necessity of larger divalent cation than Mn^{2+} , facilitating the substitution of Na for Li, i.e., the substitution of larger cation for smaller one.

Occurrence

The Dohnoiri mine is located at about N 36°32'36" and E 139°38'20", and operated for manganese ore about 50 years ago. Currently some small dumps are left around the collapsed portal. Considering from the contents of dump materials and exposures nearby the portal, the ore deposit consists of ore beds of manganese minerals developed within bedded or massive chert beds. The principal ore-forming minerals are rhodochrosite, rhodonite, hematite, braunite and natronambulite. Gangue minerals are quartz, albite and barite. Besides them minor pyrite grains are involved within low grade manganese ores composed of rhodonite and quartz.

Natronambulite occurs as light orange to brownish yellow very minute grains forming bands composed of aggregates with fine-grained quartz. Their individual grains are of microscopic dimensions (Fig. 1). Natronambulite is easily recognized due to the color. It's light orange to brownish yellow color is discernibly different from light pink color of rhodonite. Aggregates of natronambulite-quartz and rhodonite-quartz are closely accompanied by light brownish hematitequartz bedded aggregates that develop conchoidal fracture if crashed. This suggests the practical absence of recrystallization, meaning the extremely weak metamorphic effects. Some of ores accompany colorless to gray chert with conchoidal fracture as well as the ores. Also, some ores are intersected by veinlets of colorless barite, which is uniquely coarser-grained, and a light yellow fibrous amphibole developed parallel to veinlets walls attaining a centimeter long.

Under the microscope, natronambulite occurs as a minute grain with quartz and rhodochrosite, though the mineral grains tend to form minute aggregates of less than millimeter size (Fig. 2). Some interstices of quartz grains are filled in albite of similar size. In aggregates of natronambulite and quartz, veins of minute fiber crystals of amphibole are observed (Fig. 3). Although the amphibole resembles Mn-rich richterite, it has been subjected to current study. It is visually distinguishable from natronambulite due to its fibrous nature and less orange or brownish and lighter yellowish color.

Chemical composition

The electron microprobe analyses were made using Link Systems energy dispersive X-ray spectrometer (QX-2000). The results of the analyses are given in Table 1. All of them have low totals, which are considered to be due to Li_2O and H_2O . The basis of calculation taking Si=5 is warranted due to the absence of Al_2O_3 . Since all of them gave slight excess of divalent cations, a part of Ca is transferred to the site for Li and Na, requiring slight deficiency of H_2O for charge balance. The addition of above components is validated because all summations of analyses are nearly ideal.

The estimated Na/(Na+Li) ratio situated between the original (0.80) and Kombat material (0.60) serves to confirm the existence of continuous solid solution series within the compositional range of natronambulite, if the appropriate amount of CaO is involved.

Table 1. Chemical composition of natronambulite

| | 1 | 2 | 3 |
|--------------------|--------|-------|-------|
| SiO ₂ | 49.27 | 48.61 | 48.50 |
| MgÕ | 1.22 | 1.07 | 0.80 |
| MnO | 42.51 | 39.09 | 39.19 |
| CaO | 2.31 | 4.87 | 4.96 |
| Na ₂ O | 2.85 | 3.66 | 3.56 |
| Li ₂ O* | 0.86 | 0.40 | 0.47 |
| H ₂ O* | 1.36 | 1.30 | 1.32 |
| Total | 100.38 | 99.00 | 98.80 |
| Si | 5 | 5 | 5 |
| Mg | 0.185 | 0.164 | 0.123 |
| Mn | 3.654 | 3.405 | 3.422 |
| Ca | 0.251 | 0.536 | 0.548 |
| Na | 0.561 | 0.730 | 0.712 |
| Li | 0.349 | 0.165 | 0.195 |
| Н | 0.910 | 0.895 | 0.907 |

*: calculation



Fig. 1. The banded structure of natronambulite (orange), rhodonite (pink) and quartz (white). Specimen is approximately 12 cm wide.



Fig. 2. Photomicrographs of natronambulite. a: one polar, b: crossed polars. Field view is approximately 1 mm wide.

Fig. 3. Photomicrographs of minute prismatic crystals of Mn-bearing amphibole including in natronambulite. a: one polar, b: crossed polars. Field view is approximately 1 mm wide.

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| Dohnoiri | | | Tanohata | | | |
|----------|--------|----------------|---|-------|-----|--|
| dobs. | dcalc. | Ī | <u>h k l</u> | dobs. | Ī | <u>h k l</u> |
| 7.11 | 7.09 | 13 | 110 | 7.13 | 47 | 1 1 0 |
| 6.70 | 6.69 | 15 | 001 | 6.70 | 44 | $\underline{0} 0 1$ |
| | | | _ | 5.21 | 1 | 1 <u>2</u> 0, 101 |
| 4.76 | 4.75 | 9 | 111 | 4.77 | 7 | 111 _ |
| | | | | 4.15 | 1 | 021,121 |
| | | | | 3.94 | 2 | $\frac{1}{2}$ 2 0 |
| | 0.54 | - | - | 3.80 | 2 | $\frac{2}{2}$ 10 |
| 3.56 | 3.54 | 6 | 220 | 3.359 | 100 | $2 \underline{2} 0, 1 2 1$ |
| 3.35 | 3.35 | 16 | 0 0 2 | 3.348 | 40 | 131,002 |
| 2 17 | 2.17 | 20 | 102 | 3.280 | 1 | $\frac{0}{1}$ 1 2 |
| 3.17 | 3.17 | 28 | $\frac{1}{1}$ 0 2 | 3.164 | 20 | 1 0 2, 0 3 1 |
| 3.09 | 3.09 | 14 | $\frac{1}{2} \frac{1}{2} \frac{2}{0} 2 \frac{1}{2} 1$ | 3.078 | 45 | 221 |
| 3.07 | 3.07 | 17 | 2 3 0, 2 2 1 | 2 007 | 2 | 1 1 2 |
| 3.00 | 3.01 | 21 | $1 \frac{1}{2} \frac{1}{2}$ | 3.007 | 24 | $1 \frac{1}{2} \frac{2}{2}$ |
| 2.97 | 2.97 | 100 | $\frac{0}{1}$ 2 2, 1 1 2 | 2.972 | 54 | $\frac{0}{1}$ 2 2, 1 1 2 $\frac{1}{1}$ 4 0, 1 0 2 |
| 2.91 | 2.00 | 100 | 140 | 2.917 | 1 | $1 \frac{4}{2} 0, \frac{1}{1} 0 2$ |
| | | | | 2.829 | 1 | 122, 122 |
| 2 71 | 2 71 | 73 | 220 | 2.778 | 15 | 11220 |
| 2.71 | 2.71 | 13 | $2\frac{2}{4}0$ | 2.709 | 15 | 1 1 2, 2 2 0 |
| 2.01 | (2.67 | 15 | $\frac{1}{2}$ 1 2 2 2 1 | 2 618 | 15 | $\overline{2}1222\overline{1}$ |
| 2.62 | 2.61 | 11 | 141 | 21010 | 10 | , |
| 2.50 | 2.50 | 10 | $\frac{1}{3}$ 2 0 | 2.506 | 38 | $\overline{2}$ 2 2, $\overline{3}$ 2 0 |
| 2.42 | 2.42 | 14 | 122 | 2.435 | 1 | 311 |
| | | | | 2.402 | 1 | $\overline{3}$ 2 1 |
| 2.22 | (2.35 | 2 | 202 | 2.373 | 15 | $2\overline{2}2,\overline{3}30$ |
| 2.33 | 2.32 | Z | $\overline{1}$ 5 0 | | | |
| 2.23 | 2.24 | 13 | 310 | 2.245 | 1 | $1\overline{4}2,23\overline{1}$ |
| | | | _ | 2.230 | 1 | 3 <u>1</u> 0 |
| 2.22 | 2.22 | 16 | 042 | 2.221 | 6 | 013 |
| 2.20 | 2.20 | 17 | 250,103 | 2.198 | 23 | <u>3</u> 01,103 |
| | | | _ | 2.169 | 3 | 113 |
| 2.11 | 2.10 | 11 | 251 | | | . |
| 2.06 | 2.06 | 12 | 3 1 1 | 2.081 | 2 | <u>3</u> 22_ |
| | | | | 1.995 | 4 | 213,203 |
| | | | | 1.931 | 3 | |
| | | | | 1.889 | 2 | |
| | | | | 1.879 | 2 | |
| | | | | 1.859 | 6 | |
| | | | | 1.823 | 4 | |
| | | | | 1./// | 2 5 | |
| | | | | 1.705 | 1 | |
| 1 605 | 1 601 | $\gamma\gamma$ | $\overline{2}$ 4 3 | 1.740 | 2 | |
| 1.095 | (1.686 | 22 | $\frac{2}{3}$ $\frac{1}{5}$ $\frac{1}{5}$ | 1.097 | 2 | |
| 1.689 | 1 683 | 10 | 160 | | | |
| | (1.675 | | $3\frac{6}{6}1$ | 1 673 | 20 | |
| 1.674 | 1.674 | 15 | $01\frac{1}{4}$ | 1.070 | 20 | |
| | | | | | | |

Table 2. The powder X-ray diffraction pattern for natronambulite.

Dohnoiri: present study Tanohata: Matsubara *et al.* (1985)

X-ray crystallography

The powder X-ray diffraction pattern was obtained using a Gandolfi camera, 114.6 mm in diameter, employing Cr $K\alpha$ radiation (Table 2). The unit cell parameters are: a=7.624(5), b=11.605(9), c=6.738(3) Å, $\alpha=92.96(6)^{\circ}$, $\beta=94.75(5)^{\circ}$, $\gamma=106.26(6)^{\circ}$, V=568.7(6) Å³. Though the pattern is apparently similar to that of rhodonite or pyroxmangite, it corresponds well to that of the original one (Matsubara *et al.*, 1985). As the compositional variation of natronambulite yields minor variation in powder Xray diffraction pattern, it is very likely that the similarity to rhodonite or pyroxmangite has induced the confusion with this mineral.

Consideration

In addition to the retained bedded texture, fine grain size of natronambulite as well as the associated quartz suggests the lesser degree of influence of metamorphism. It is highly likely that all the component of natronambulite had been concentrated prior to its formation.

The Na: Li ratio in the original natronambulite is 0.82:0.18. Matsubara *et al.* (1985) have considered that the entrance of Na in place of Li necessitates that of larger cation in place of Mn such as Ca on the basis of the unique original analysis in which the quantitative determination was made by wet method. The chemical analyses of the present material substantiate this conclusion; higher CaO content than the original one and no mutual correlation between Li_2O and CaO contents.

Natronambulite and nambulite are apparently characterized by orange to brownish yellow color. However, they have been miss-identified as rhodonite or pyroxmangite to date, since some of high iron-bearing rhodonite has orange brown color similar to those of natronambulite and nambulite. This evidence may lead to the new find of these Li-bearing pyroxenoids in lower grade siliceous manganese ores from weakly to highly metamorphosed bedded manganese deposits.

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